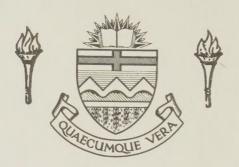
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THE UNIVERSITY OF ALBERTA PALSAS IN BAKER CREEK BASIN N.W.T.: AN ECOSYSTEMATIC STUDY

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A THESIS

SUBMITTED TO THE FACULTY OF GRADUATE STUDIES AND RESEARCH
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ABSTRACT

Peatland (both minerotrophic and ombrotrophic) extends from southern Canada northward to include parts of the continuous permafrost zone. In the subarctic, discontinuous permafrost is restricted mainly to peatlands and often occurs as raised peat landforms termed palsas. Palsa uplift results from water migration to the freezing plane as well as buoyancy and water volume expansion upon freezing. The genesis of palsas (that form in minerotrophic topogenous carr peatlands) occurs in Carex - Tomenthypnum meadows where individual peat hummocks develop a perennially frozen core. Organic accretion and permafrost aggradation cause these hummocks to become larger and more numerous.

Surface peat of the larger hummocks is drier than surface peat of the Carex - Tomenthypnum hummocks. As a result, the Betula - Pohlia association becomes dominant. Over the years, permafrost aggradation coupled with slow peat infilling of depressions between hummocks results in a raised peat landform having a perennially frozen core.

Termed a palsa, the uplifted peat surface is drier than surface peat of the <code>Betula - Pohlia</code> hummocks. The dryness results in the dominance of the <code>Ledum - Cladonia</code> association as well as the climatic climax <code>Picea - Cladonia</code> association.

Environmental differences exist between plant associations. However, within each association the environment is uniform. Differing plant associations provide habitats of varying suitability for small mammals. Habitat suitability is greatest in the Betula - Pohlia plant association and decreases toward the climatic climax Picea - Cladonia association.

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ACKNOWLEDGEMENTS

The research and writing of this thesis owes its culmination to my parents. Through the years their encouragement and sacrifice have made my education possible.

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Finally, I mention Roscoe Rawlins, whose inspiration during the final drafts re-established the goals of this work.



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CHAPTER ONE

INTRODUCTION

1.1 INTRODUCTION

Permafrost refers to the thermal conditions of earth materials when their temperature remains below 0°C continuously for a number of years. It occurs where there is a negative heat balance and is defined exclusively by temperature. At the southern fringe of the subarctic, permafrost is restricted mainly to peatlands (Brown 1970) and often occurs in distinct raised peat landforms termed palsas. Palsas generally are less than 100 m in diameter and vary in height from 1 - 3 m (Sjörs 1961).

Palsas have attracted the interests of researchers in Europe (Rapp and Rudberg 1964, Mörnsjö 1968, Ruuhijarvi 1968, Salmi 1968, Schenk 1968, Svensson 1968, Lundquist 1969, Mörnsjö 1971, Wramner 1972, Warmner 1973, Sollid and Sorbel 1974, Vorren and Vorren 1975) and in North America (Sjörs 1959, Railton and Sparling 1973, Zoltai and Tarnocai 1971, Zoltai 1972, Thie 1974, Brown 1977, Reid 1977). Until the present study, however, no work has been available that analyzed palsa development within the Precambrian Shield of Canada or similar physiographic regions elsewhere.

1.2 PURPOSE OF RESEARCH

This study analyzes the interrelations of vegetation and physical environment that form palsas in a small area of the Canada Precambrian Shield. Palsa development is linked to plant succession, therefore seral



trends within the study area are examined. Individual successional associations function as habitat for small mammals, thus the influence of palsa formation on habitat development is analyzed. To understand these interactions I studied the geomorphology, microclimate, plant succession and small mammal distribution of three separate depressions near Yellowknife, Northwest Territories.

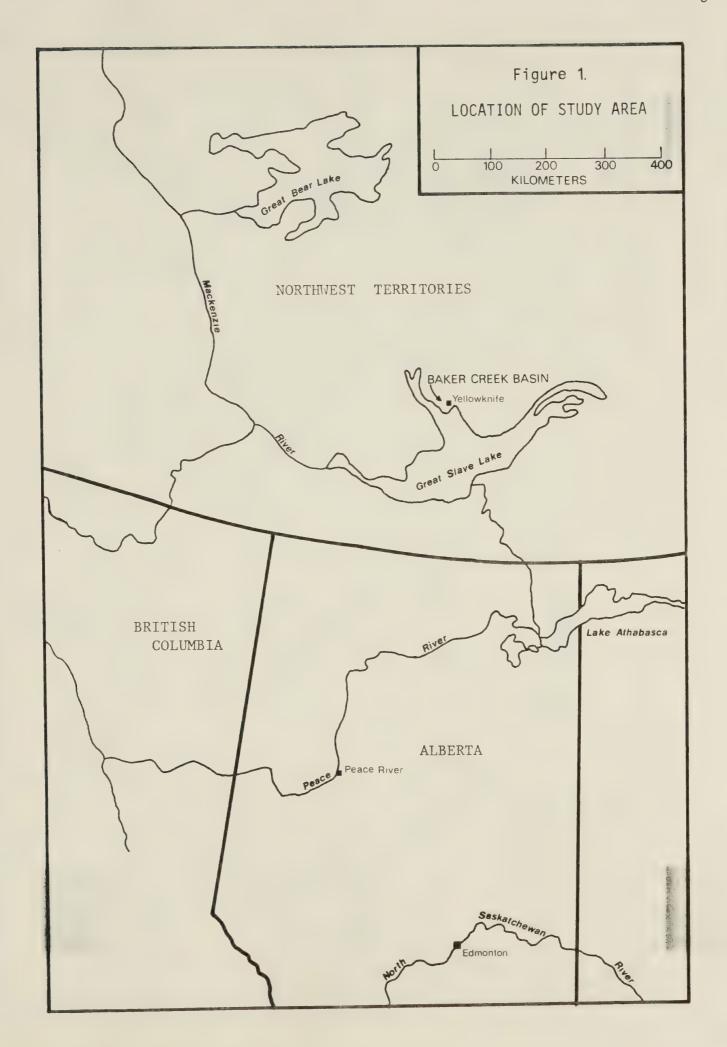
1.3 ECOSYSTEMATIC APPROACH

The ecosystematic approach is used in the present study to analyze and classify varying combinations of the physical environment as they interact with the flora of palsas to form small mammal habitat. The ecosystem concept has been used by many researchers to study the interaction of biotic communities with the physical environment (Stoddart 1965, Odum 1969, Gill 1972). Vegetation units formed by the plant-habitat complex are defined as plant associations. The association concept implies a uniform aggregation of plants which are in dynamic equilibrium with the physical environment. I used this concept to analyze the flora of the study area.

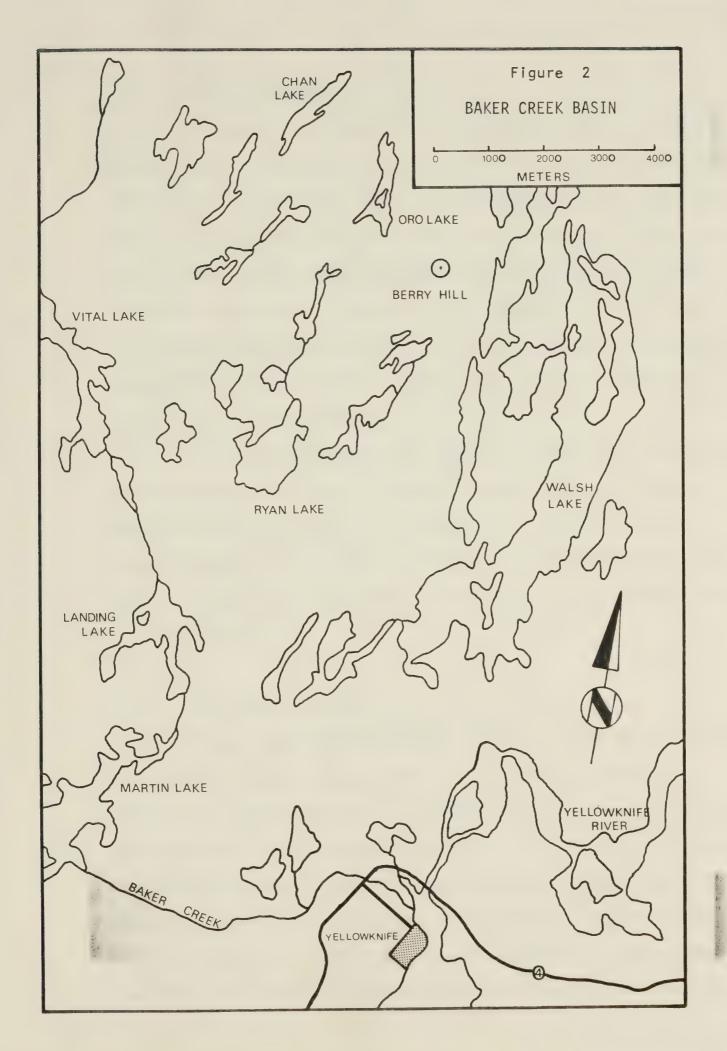
1.4 STUDY AREA

Research for the present study took place in the northeast sector of Baker Creek Basin, situated on the west side of Yellowknife Bay, between 62° 28' and 62° 43' north and 114° 17' to 114° 31' west (Figure 1). Intensive work was conducted between Oro and Chan Lakes (Figure 2).











1.4.1 STUDY AREA CLIMATE

Because of the study area's location, continentality plays a major role in its climate. Yellowknife is situated approximately 565 km from the Arctic Ocean and 1210 km from the Pacific. The Rocky Mountains accentuate the great distance from the Pacific by impeding flow of maritime air toward the interior of the continent. Climatically, the study area is in the subarctic (Köppen and Geiger 1953). Precipitation and temperature regimes of the study area are represented by meteorological data for Yellowknife (15 km distant) covering the period 1941 to 1970 (Figure 3).

The continentality of the climate is shown by a large annual temperature range (Figure 3). Summers are cool, with July having a mean temperature of 15.6°C. Winters are cold, having an average January temperature of -28.3°C. Yellowknife's precipitation is moderately concentrated in late summer and early fall, and averages 250 mm annually (Figure 3).

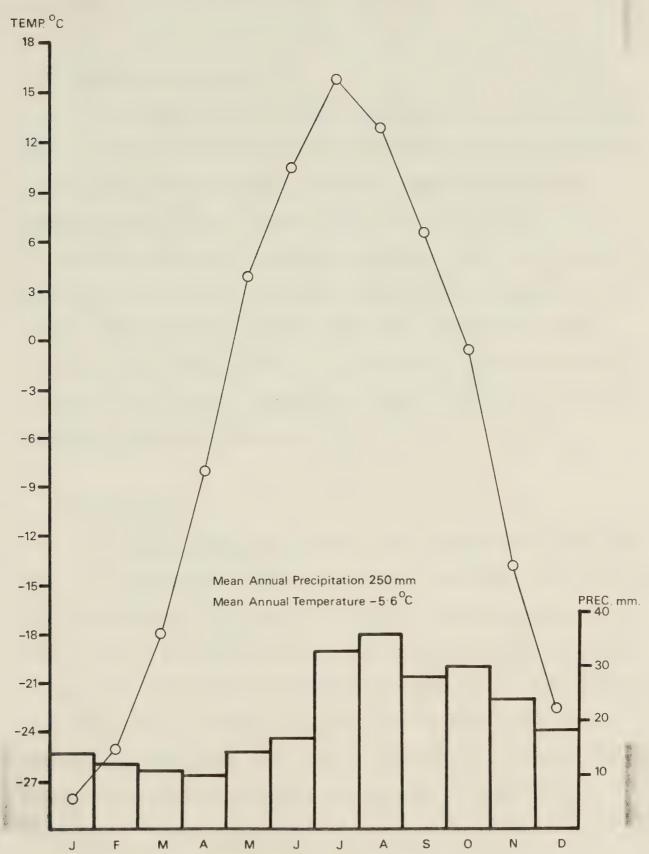
1.4.2 STUDY AREA PHYSIOGRAPHY

Baker Creek Basin is situated on the Canadian Shield. Eightyfive per cent of the basin is underlain by Archaean granodiorite,
granite and allied intrusive rocks (Wight 1973). The remaining 15 percent consists of the Yellowknife Greenstone Belt. Wight (1973) suggested
that the region east of the basin was deglaciated 7000 years BP.
Berry Hill which is the highest point in the basin is approximately
275 m above sea level. The lowest point in the basin is 167 m above
sea level. Most aspects of relief are related to structure and lith-



Figure 3

Mean Monthly Temperature and Precipitation:
Yellowknife Northwest Territories





ology, although glaciation has imposed its characteristic surface morphology.

1.4.3 STUDY AREA VEGETATION

The study area lies in the open Boreal Woodland (Rowe 1959). Plant associations of the basin are divided into two major categories (Landals 1970), those occupying the upland ridge sites and those colonizing major valleys or depressions. The upland sites contain six distinct associations, while the depressions (the locations most important to the present study) are occupied by four separable associations. Dominant species of the study area depressions include:

Carex aquatilis, Betula pumila var. glandulifera, Ledum palustre var. decumbens, Picea mariana, Tomenthypnum nitens, Pohlia nutans, Cladonia coccifera, and Cladonia nivalis.

1.5 FIELD SEASON

Field reconnaissance of Baker Creek Basin began 1 July 1973 when I was a research assistant to a Master's candidate in the Department of Geography, The University of Alberta. We used the Curtis (1959) transect sampling method to study vegetation in selected areas of the basin. That work was completed 31 August 1973. It was during that period that I selected the present area of study. The 1974 summer field season began on 1 June and extended to 31 August. Winter snow and temperature observations were made at the study site in March 1975. I spent a total of five months in the field.



1.6 CHOICE OF STUDY AREA

I chose the northeast sector of Baker Creek Basin (Figure 2) as the area of study for three reasons: the city of Yellowknife is relatively close to the study site, thus logistic support was reasonably available; the study location is situated northwest (generally upwind) of the mining operations of Yellowknife and far enough away (15 km) that air pollution from those operations has little or no environmental effect; and based on my familiarity with the location from having worked there during July and August 1973, I found the conditions well suited for this study.



CHAPTER TWO

VEGETATION

2.1 INTRODUCTION

Vast areas of Canada are covered by deposits of peat. Brown and Williams (1972) estimated that this type of terrain covers approximately 1,300,000 km². It extends from southern Canada northward to include parts of the continuous permafrost zone. This unique terrain is variously referred to as bogland, muskeg, organic terrain and peatland. The term peatland has many modern connotations, but for engineering purposes it is defined as terrain composed of a living organic mat of mosses, sedges, and/or grasses, with or without tree growth, and underlain by a highly compressible mixture of partially decomposed and disintegrated organic material commonly known as peat (MacFarlane 1959).

Organic terrain is composed of vegetal remains and its preserved state as peat reflects ordered processes (successional phenomena) which are a function of climatic, edaphic and biotic influences. Peatland can be classified by three methods: stratigraphically, geochemically and hydrotopographically.

Stratigraphically, peatlands rest directly on mineral ground or are underlain by limnic, organic sediments (Mörnsjö 1971). Peatland which is directly underlain by mineral ground indicates that the peat formed by paludification of the land. Peatland underlain by limnic,



organic sediments indicates that peat formation was initiated through the shoaling of a lake by an input of detritus. This process is called terrestrialization. Based on these broad stratigraphic characteristics, peatlands are genetically classed as paludificational peatlands or terrestrializational peatlands (Mörnsjö 1971).

The origin of water and concomitant nutrient supply is a primary factor differentiating peatland types. All peatlands receive water directly from precipitation; in some types of peatland, precipitation is the only source of moisture and consequently the primary source of mineral salts (Sjörs 1959). Such peatlands are ombrotrophic* and are sometimes called bogs. Geochemically, ombrotrophic peatlands have a pronounced deficiency in mineral nutrients, have highly acidic water and peat and low Ca/Mg ratios. Ombrotrophic peatlands are composed of peat types largely from *Sphagnum* spp. remains and physiognomically are of the domed type (Mörnsjö 1971).

Minerotrophic peatlands, or fens, receive varying amounts of water which have been in contact with mineral ground either as ground water or as surface runoff as well as precipitation. Geochemically, minerotrophic peatland has a high content of mineral nutrients and a higher Ca/Mg ratio than ombrotrophic peatland. Minerotrophic influence in peatlands usually results in a richer, more varied vegetation composed of species normally not present in ombrotrophic areas.

^{*} Appendix 1 defines relevant terms



Hydrotopographically, ombrotrophic peatland is separated into three types; eccentrically domed bog peatland, concentrically domed bog peatland and forested bog peatland (Mörnsjö 1971). Eccentrically domed bog peatland is developed on gentle slopes of permeable ground in areas of high precipitation. From the top of its dome (located at the upslope margin) the peatland surface slopes downward as does the ombrogenous ground water level (Mörnsjö 1971). Developed in areas of flat terrain, concentrically domed bog peatland represents the classic raised bog with a centrally located dome. The ombrogenous ground water level slopes gently in all directions.

Ground water level in both the eccentrically and concentrically domed bog peatlands is high throughout the year. These peatlands are unwooded. Hydrologically the forest bog peatland is characterized by oscillations of ground water. From autumn to spring the ground water level is close to the peat surface. However, a pronounced water-deficit in summer favours the growth of trees. The surface of this peatland is only slightly domed, usually in a centrical manner.

Hydrotopographically, minerotrophic peatlands are classified into two separable types. The first type is termed rheo-geogenous, while the second is called topogenous (Mörnsjö 1971). Topogenous peatlands are separated into two subtypes termed topogenous-fens and topogenous-carrs. Rheo-geogenous minerotrophic peatlands are conditioned by laterally mobile ground water of continuous flow. These peatlands have a sloping surface and are usually found in areas of high precipi-



tation. Topogenous minerotrophic peatlands are developed in basins that contain stagnant ground water. These peatlands have a level surface. Topogenous fen peatlands maintain a high ground water level throughout the year. Topogenous carr peatlands are influenced by oscillating ground water levels. Primarily, the topogenous carr peatlands are composed of wooded peat whereas moss peat is the primary component of topogenous fen peatlands. It would seem (based on Mörnsjö's (1971) definitions) therefore, that the present work was conducted in a minerotrophic topogenous carr peatland.

Previous botanical research in the vicinity of the north arm of Great Slave Lake is limited (Landals 1970, Wight 1973). Early work (Richardson 1851, Macoun 1877, Preble 1908) was conducted along the margins of major waterways with the consequence that inland areas were unknown. With the advent of air travel that pattern changed. Raup (1946), Cody and Chillcott (1955), Scotter (1966) and Larsen (1971) conducted research proximal to the east arm of Great Slave Lake. Research south of the lake was executed by Harper (1931), Raup (1946) and Cody (1956). In the west, botanical investigations were made by Raup (1947), Cody (1960) and Theiret (1961).

With the exception of Raup's (1946) and Larsen's (1965, 1971) brief statements, no work is available for the Great Slave Lake area that analyzes plant succession in peatland areas. Therefore the present study is a contribution to the ecological knowledge of the Great Slave Lake area.



2.2 PROCEDURE

Phytosociologic methods as developed by Braun-Blanquet (1932) and modified by Gill (1971) are applied to the analysis of vegetation in the study area. The Braun-Blanquet method of community analysis is described as follows: a survey is made of a homogeneous plant community and all plant species present are recorded. Within the uniform area, preferably at a nodal position, a representative sample plot is established. If most of the listed species are not located in the sample plot it is enlarged until all or nearly all species are encountered. Once the minimum necessary plot size is established, floristic and environmental analyses are conducted.

The Braun-Blanquet method of community analysis is criticized (Janssen 1967) because the investigator selects homogeneous sample plots and does not take into consideration transitional vegetation between communities or associations. However, plant associations of the present study area have boundaries that are distinct with little ecotonation between them. Therefore, the Braun-Blanquet method of community analysis is an adequate study tool. Vegetation distribution in the study area was mapped from the ground with the aid of aerial photographs and from the air in a small aircraft.

2.2.1 SIZE OF SAMPLE PLOT

Sample plot size is selected to ensure that plot areas are sufficiently representative of each association. The problem in determining plot size is one of keeping it sufficiently small to handle while ensuring that it is large enough to contain representative



association information. In the present research area, the restricted distribution of some associations made it necessary for certain plots to be elongated. Plot sizes ranged from 4 m² in the smaller *Carex* associations to a maximum of 100 m^2 in the larger Picea stands. Twentyfour such plots are analyzed (six per association).

2.2.2 ANALYSIS OF SAMPLE PLOTS

TABLE 1

After a tentative vegetation type is determined for an association, stratification of the component species in the sample plot is assigned as shown in Table 1.

STRATIFICATION INDEX

А	Tree Layer	above 6 m
В	Low Tree/Tall Shrub Layer	below 6 m
С	Shrub-Herb Layer	below 1 m

Moss/Lichen Layer D_1 Moss/Lichen Layer on wood D_2

A number (Table 2) denoting sociability is given to each species in the sample plot.

on peat

TABLE 2. SOCIABILITY INDEX

- Growing singly 1
- Grouped or tufted 2
- Growing in small patches or cushions
- Growing in extensive patches or carpets 4
- 5 Forming pure populations



A second number (Table 3) denoting average vigor of the species being evaluated is added.

TABLE 3.	VIGOR INDEX
1	Dead
2	Dying
3	Poor
4	Good
5	Excellent

A figure from the cover-abundance index (Table 4) is given to each species.

TABLE 4.	COVER-ABUNDANCE	INDEX
1/1066 7 .		- INDLA

- x Sparsely present, cover very small
- 1 Plentiful, covering 5 percent of area
- 2 Numerous, covering 6 25 percent of area
- 3 Covering 26 50 percent of area
- 4 Covering 51 75 percent of area
- 5 Covering more than 75 percent of area

Species importance in the study plot is assigned as shown in Table 5.

TABLE 5. SPECIES IMPORTANCE

- 1 0 19 percent of total plants
- 2 20 39 percent of total plants
- 3 40 59 percent of total plants
- 4 60 79 percent of total plants
- 5 80 100 percent of total plants



Using a Swedish increment borer, trees were aged by taking increment cores from five trees in each plot. Annual rings were counted by binocular microscope employing methods used by Parker (1971). The stems of Salix in some plots were cut and a direct growth-ring count was made in the field with the aid of a hand lens.

Once floristic analysis was complete, voucher specimens of each species of plant were collected from the sample plots. The specimens were processed and returned to The University of Alberta where final identification was made. Voucher specimens were deposited in the Herbarium of the Botany Department, The University of Alberta.

2.3 PLANT ASSOCIATIONS

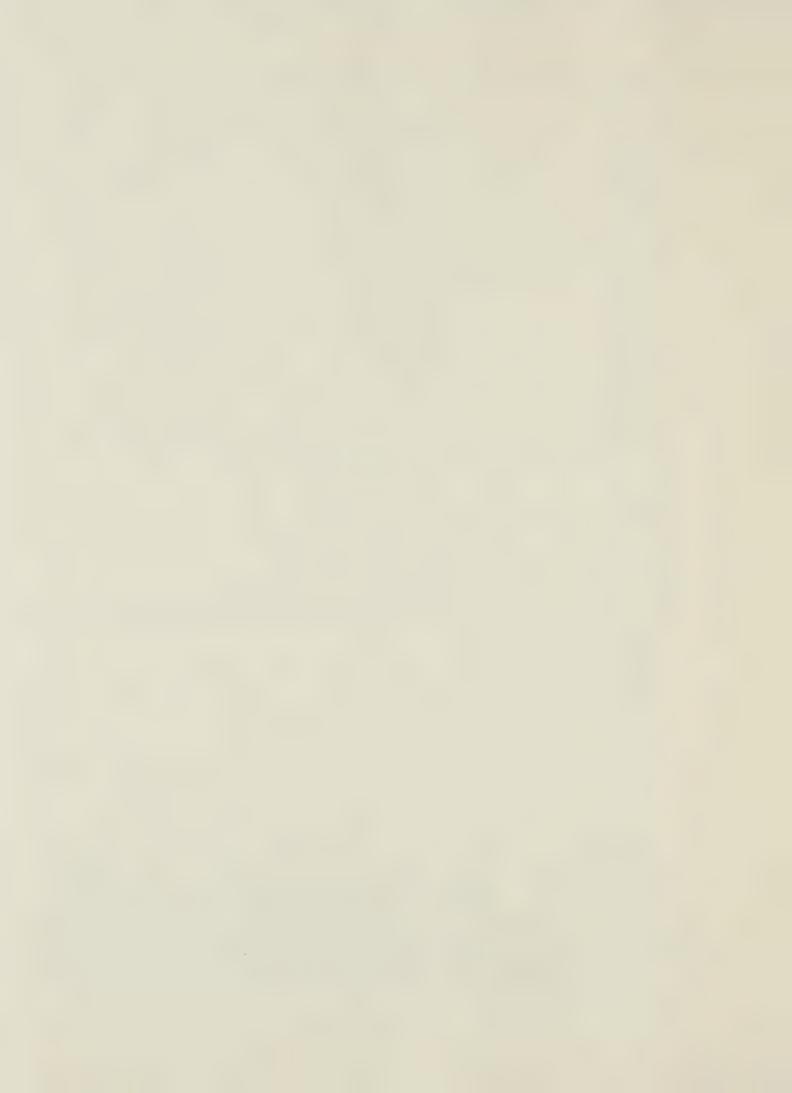
Tables are often used to simplify the study of vegetation associations. The Associations Table (Table 6) depicts the cover abundance of species occuring in each study plot. Major inflection points are empirically seen between the Carex - Tomenthypnum and Betula - Pohlia associations. Comparison of active layer depths indicates relative homogeneity within each association but differences between them.

Tables 7 - 10 present separately, in tabular form, each association of the study area. The association tables denote the stratification, sociability, vigor, cover - abundance and importance values (Tables 1 - 5) for each species of a particular association. These indices are followed by their averages. Indicator species for each association are determined from those which have a high importance value and a high index of cover - abundance. Two indicator



ASSOCIATIONS TABLE

ASSOCIATION	Care	x - 7	Carex - Tomenthypnum	thyk	mum			Betui	Betula - Politia	hlia			7	Lec'um - Cladonia	- Clad	eiuc			0	Picea - Cladonia	Clado	nia		
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SPECIES																								
Tomenthypnum nitens Carex aquatilis Carex rostrata Myrica gale Salix glauca Betuia pumila var. glandulifera Potentilla palustris Eriophorum angustifolium Smilacina trifolia	n m m	omm ← ×	× + + ×	72 X X X X X X X X X X X X X X X X X X X	₩ W W W W W W W W W W W W W W W W W W W	- N NX -		ω ← α ← α ← ×	× 0 0 × ×		- N N													
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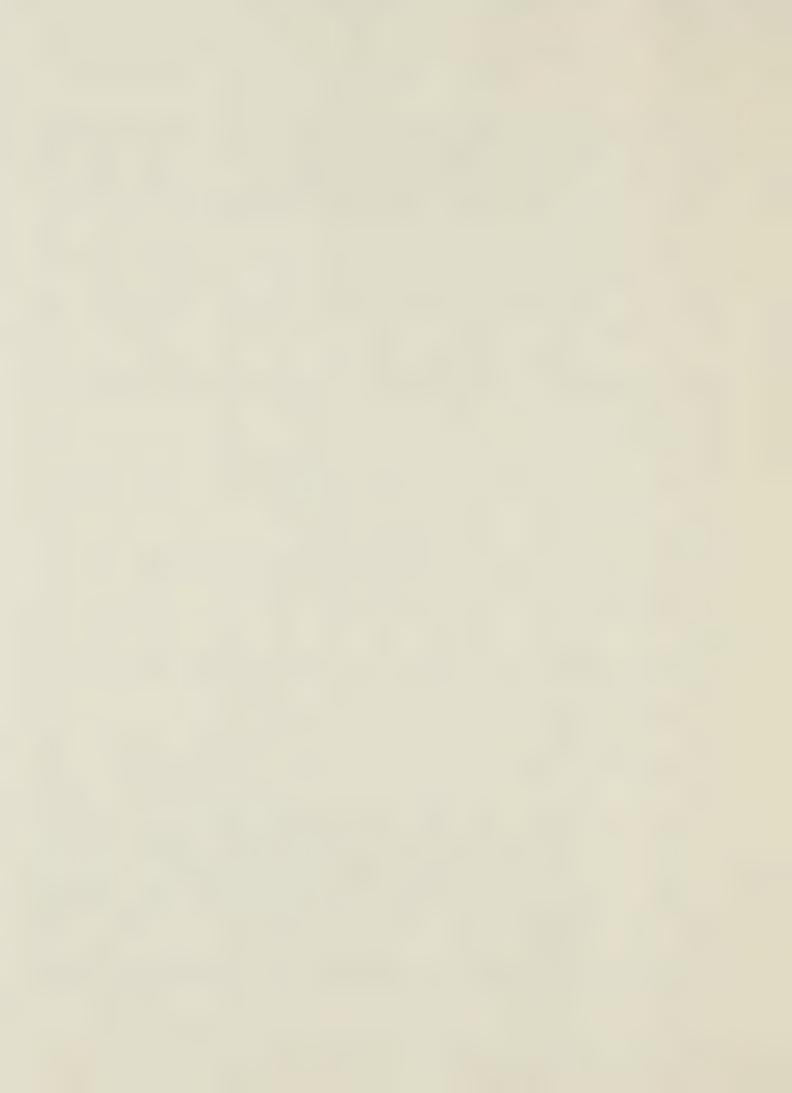
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		Plot Number		Species	Carex aquatilis	Carex rostrata	Myrica gale	Salix glauca	Betula pumila var. glandulifera	Potentilla palustris	Eriophorum angustifolium	Smilacina trifolia	Tomenthypnum nitens
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TABLE 7



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	Plot number	Species	Picea mariana	Picea mariana	Ledum palustre var. decumbens	Vaccinium vitis - idaea	Empetrum nigrum	Picea mariana	Rubus chamaemorus	Andromeda polifolia	Chamaedaphne calyculata	Ledum groenlandicum	Equisetum pratonse	Equisetum silvaticum	Calamagrostis canadensis	Pinus banksiana	Vaccinium uliginosum	Cladonia nivalis	Cladina mitis	Cladonia uncialis	Cladonia coccifera	Pontia rutans	Hypogymnia physodes	Cetraria pinstria
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species, one from the A, B, or C stratum and one from the D stratum (Table 1) are chosen as nomenclature.

Thus the present research was conducted within four plant associations: the Carex aquatilis - Tomenthypnum nitens, Betula pumila var. glandulifera - Pohlia nutans, Ledum palustre var. decumbens - Cladonia coccifera and the Picea - mariana - Cladonia nivalis. Plants in each association (Tables 7 - 10) are ranked according to stratum and their average cover-abundance; where two or more plants have the same cover-abundance they are ranked according to their average species importance. If two species have identical importance and cover - abundance values they are ranked alphabetically.

2.4 PRESENT PLANT SUCCESSION

There is lack of agreement in the theoretical formulation of association interactions or of successional trends (Johnson and Valentine 1971). Braun-Blanquet (1932) stated that the exchange (appearance/disappearance) of species or of changes in the quantitative relations of species continuously present in vegetation associations constitutes the succession of one association to another. Drury and Nisbet (1973) explained that the term succession is usually used to imply sequences of time. They suggested that successional series are recognized as extending from associations of low stature, few species and simple structure to associations of tall plants, many species and complex structure. This concept is analogous to Braun-Blanquet's (1932) theory of pioneer, transitional and terminal stages of succession.



The Carex - Tomenthypram association is the most simple in the present study area (Table 6). This association contains seven species in the richest sample plot. During spring and early summer, areas occupied by this association (Figures 4,5) are flooded to a shallow depth by snowmelt runoff. With the exception of permafrost lenses under hummocks, frozen material was absent from areas occupied by this association by 23 July (Figure 6, Appendix 2).

The Betula - Pohlia association occupies areas (Figures 4,5) characterized by individual peat hummocks. Eighteen species were found in the richest sample plot of this association. On 27 August no permafrost was found between the hummocks but average depth to frost beneath them was 64.0 + .87 cm (Figure 6, Appendix 2).

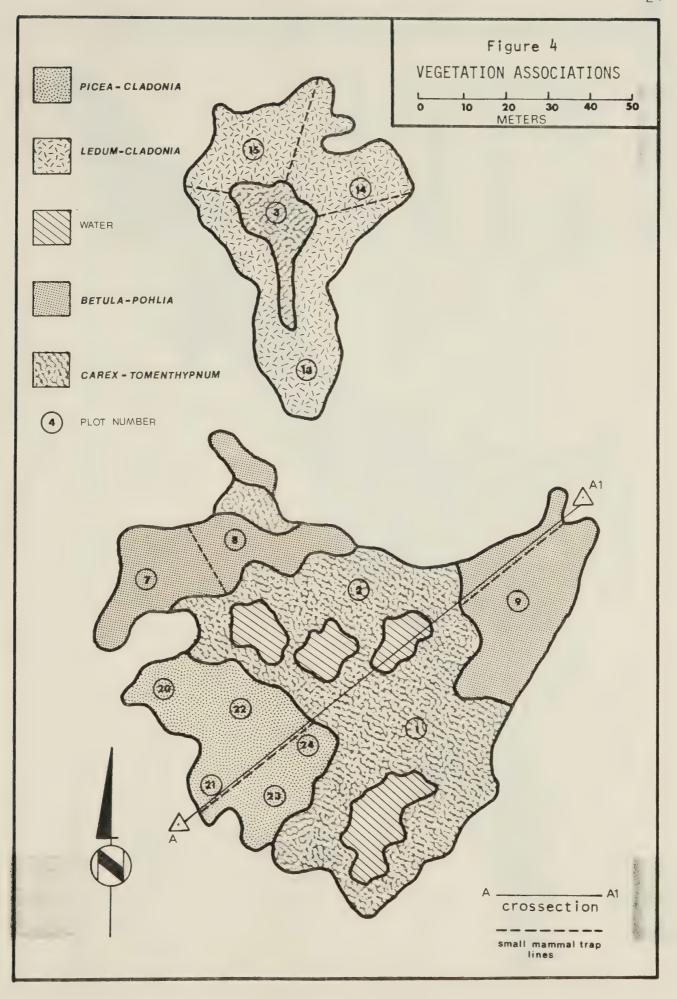
Within the Ledum - Cladonia association a maximum of 17 species were found in any one sample plot. In areas occupied by this association (Figures 4,5) individual peat hummocks are no longer discernible. On 27 August permafrost was found beneath the peat surface of this association at a mean depth of $46.1 \pm .75$ cm (Figure 6, Appendix 2).

A maximum of 17 species were found in the richest sample plot of the Picea - Cladonia association. Permafrost was found beneath the peat surface at a mean depth of 45.8 ± 1.85 cm (Figure 6, Appendix 2).

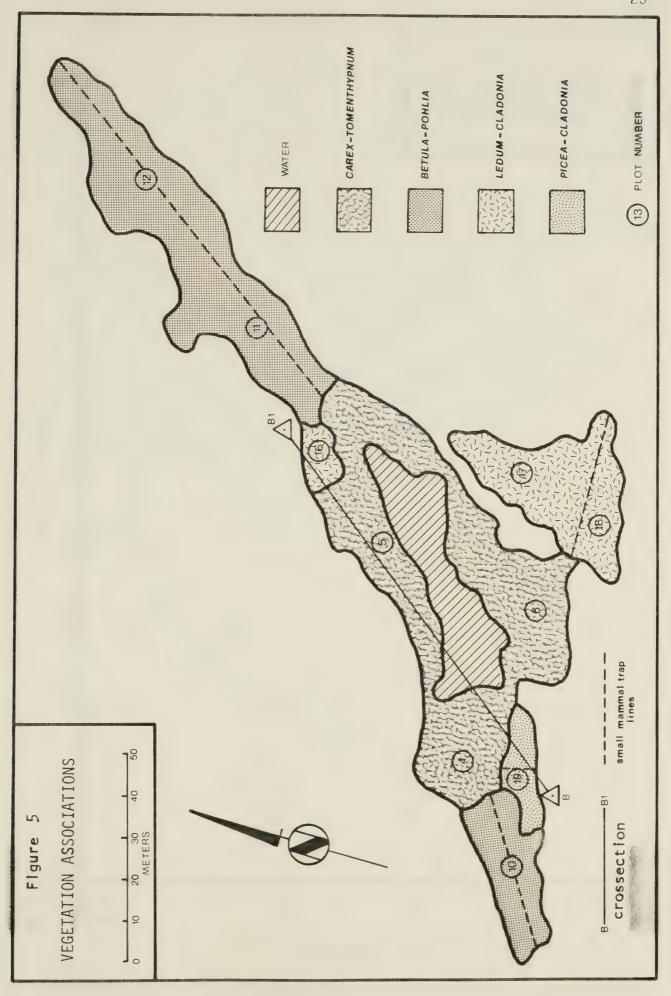
2.5 SUMMARY

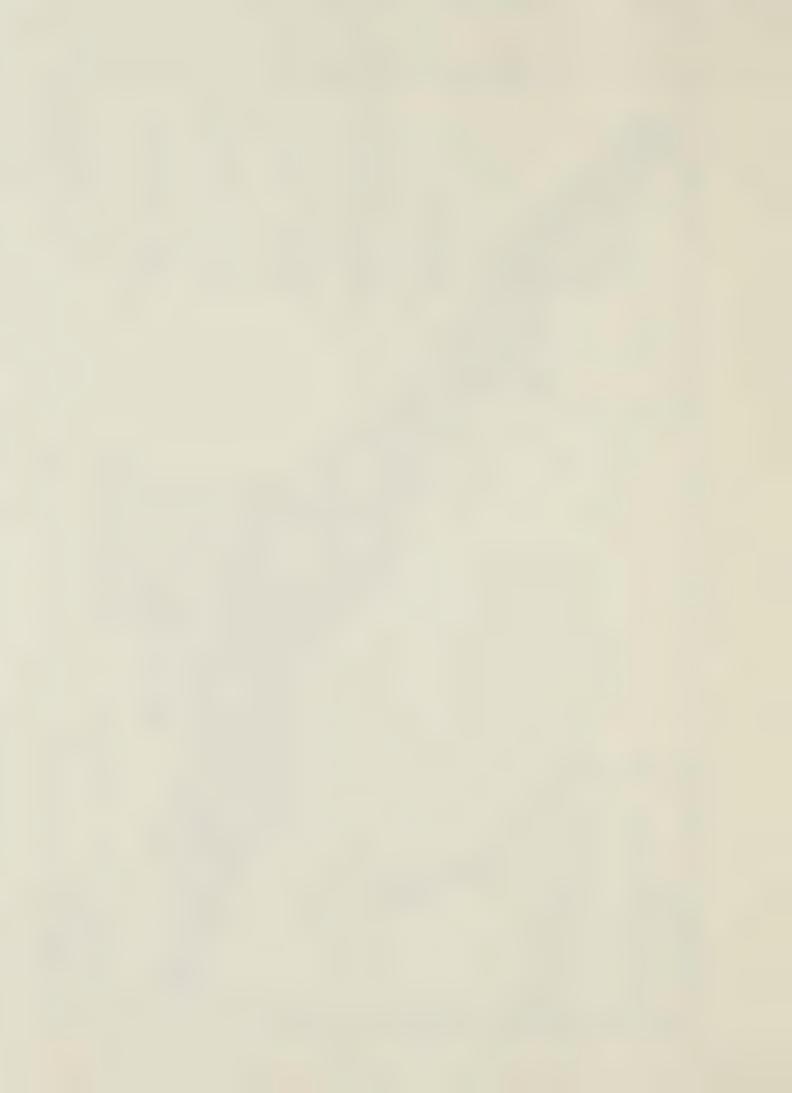
Four plant associations are present in depressions of the study area. Individually, the associations represent stages (pioneer, transitional and climax) of plant succession in a minerotrophic topogenous carr peatland.

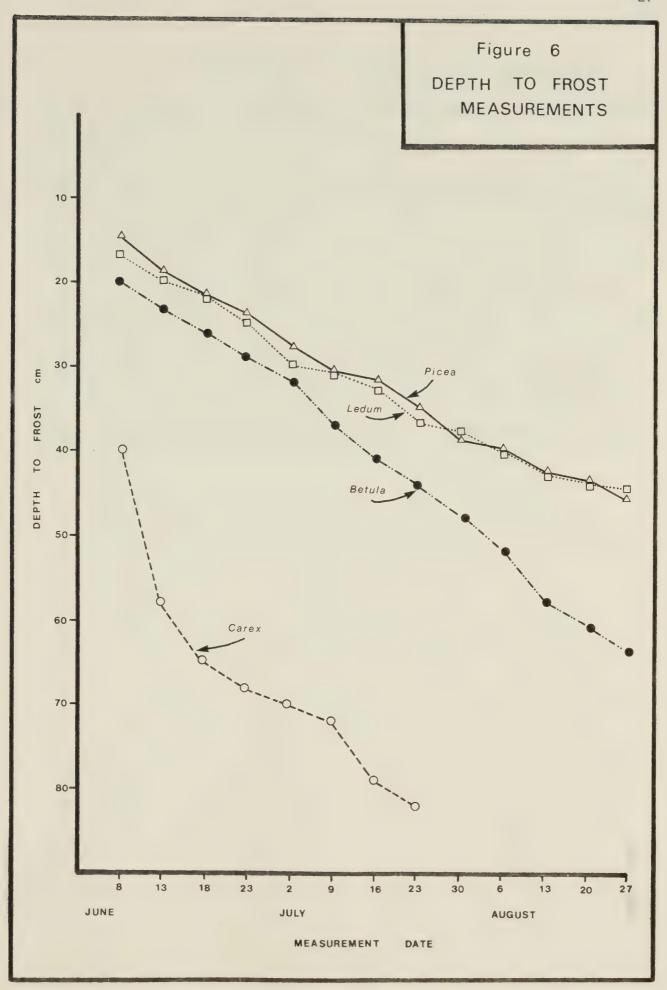














The number of transition stages in sere development (a series of sequential stages which follow one another regularly) depends upon prevailing environmental conditions as well as upon the number of species important to the seral development (Braun-Blanquet 1932). Two transitional stages (represented by the *Betula - Pohlia* and the *Ledum - Cladonia* plant associations) are noted in the present study.

The *Picea - Cladonia* association represents the climatic climax stage of succession in the depressions of the study area. Braun-Blanquet (1932) stated that a climatic climax will not be reached at the same time over an entire area. The present study confirms this, as all stages of sere development occupy the depressions.



CHAPTER THREE

PALSA DEVELOPMENT

3.1 INTRODUCTION

Palsas are raised peat landforms with a permafrost core that commonly contains ice lenses no thicker than 2 - 3 cm (Lundqvist 1969). According to Lundqvist (1969) palsa uplift may be caused by the development of such ice lenses. Mackay (1965) described peat mounds domed by pressure from methane gas in the Mackenzie Delta area and suggested (Mackay 1972, personal communication in Washburn 1973) that such doming may initiate palsa development. Similar doming of peatland and ensuing palsa development in the Rock Lake Delta area of Alberta has been attributed to methane gas pressure (Gill, unpublished data, 1976). More commonly, however, the doming of peatland and subsequent palsa development is initiated by the volume expansion of water upon freezing (Lundqvist 1969).

There is controversy within the literature concerning whether the majority of palsa uplift occurs in mineral soil below organic material, or within the peat itself. Rapp and Rudberg (1964) stated that in northern Norway there have been many observations made of palsas both with and without a core of frozen mineral soil. In northern Sweden, Forsgren (1966) found that frozen mineral soil was only a small part of the total core of permafrost in most palsas. Ice lenses were more developed in the peat than in the mineral soil below and consequently most of the palsa height was ascribed to frost-heaving



within the peat. However, Fries (1964, in Forsgren 1968) found that peat thickness in palsas was the same as in the surrounding bogs and therefore attributed most of the organic material uplift to frost-heaving in mineral soil.

Salmi (1968) proposed that during their initial stages, palsas float on unfrozen saturated peat. He stated that after initiation, the permafrost core of a palsa is supplied with water from unfrozen saturated peat and mineral soil below the permafrost layer. At this stage of uplift, the volume of a palsa above the surrounding unfrozen organic material is similar to the ratio of ice above water to mass of ice below water in a floating iceberg. Zoltai (1972) stated that buoyancy alone does not account for the total elevation of a palsa above an unfrozen fen. In addition to buoyancy release, the volume of saturated peat increases, by about nine percent upon freezing, with most expansion taking place upward in the direction of heat loss. Therefore, it would seem that a combination of water migration to the freezing plane, buoyancy and water volume expansion upon freezing is needed to explain palsa uplift.

Salmi (1968) hypothesized that palsa uplift occurs when the permafrost layer within peat attains such a thickness that its lower surface comes in contact with mineral soil; at that point its upper surface begins to rise. Because of the ubiquitousness and greater thickness of ice lenses in fine grained mineral soil than in overlying peat, he claimed that mineral soil would undergo a greater volume change



than would the organic material. Research conducted by Zoltai and Tarnocai (1971) agreed with Salmi's (1968) explanation. They stated that in northern Manitoba mineral soil beneath peat was nearly level, except under areas that were frozen, at which point there was mineral soil uplift. The authors attributed doming of the peatland to this rise in the mineral soil substratum. It would seem therefore that to fully account for organic uplift by freezing, the best explanation involves a combination of the above processes.

3.2 PALSA AND PEAT PLATEAU DIFFERENTIATION

Zoltai (1972) suggested that permafrost in peat plateaux is entirely within the peat, but in palsas, although covered by peat, permafrost extends into underlying mineral soil. Peat plateaux therefore have frozen cores, and float on a saturated peat medium, whereas palsas are firmly anchored to mineral soil by their frozen cores. Zoltai (1972) also used differences in height to differentiate between palsas and peat plateaux. He stated that palsas vary in height between 100 cm and 300 cm but may reach 500 cm. Peat plateaux on the other hand seldom exceed 120 cm. He stated that freezing of mineral soil beneath peat gives the palsa increased height compared to the peat plateau.

Comparing Zoltai's (1972) definitions of palsas and peat plateaux to statements by Rapp and Rudberg (1964), Forsgren (1966) and Salmi (1968) which indicated that organic material uplift results from permafrost aggradation within peat, it appears that a peat plateau is a young palsa whose permafrost core has not reached mineral soil. Brown



(1968) has also stated that palsas and peat plateaux of subarctic Canada are morphological variations of the same process. Therefore a differentiation in formative processes between palsas and peat plateaux may not be needed, although a differentiation in the temporal attributes of the processes might be useful.

If a differentiation between palsas and peat plateaux must be made, it should be made on a more pertinent basis than Zoltai's (1972) classification. Such a distinction can be made using surface area. Sjörs (1961) stated that palsas are generally less than 100 m in diameter and peat plateaux are greater than 100 m. I feel that this criterion (if in fact needed) should be the basis on which to distinguish palsas from peat plateaux.

3.3 PALSA DEVELOPMENT

Zoltai's (1972) work shows that vegetation plays an important role in initiating palsa development. He stated that thin layers of permafrost develop under dense clumps of *Picea mariana* growing on a fen. Because the branches retain much of the winter snowfall as qali (Pruitt 1970), depths in the snow shadow beneath the trees (qaminiq) are thinner. This allows frost to penetrate deeper into organic material beneath the trees than in surrounding areas. In summer, shade from the same trees reduces solar radiation at the ground surface and retards thawing. The resultant negative heat balance over the years causes one to several thin layers of permafrost to develop, which may coalesce into larger layers (Zoltai 1972). Permafrost aggradation continues, together



with water migration along a thermal gradient (Hoekstra 1969), until the permafrost structure is at an environmental equilibrium (Zoltai 1972). Uplift of the frozen organic material takes place because of factors described previously. Due to uplift, surface peat is better drained and becomes drier, therefore its insulating effect is sufficient to reduce heat exchange above the permafrost during the summer.

Zoltai (1972) developed field criteria for the recognition of developmental stages of palsas. Without implying time periods, he used the terms *incipient*, *young*, *mature* and *overmature* to denote these changes. Although these stages of development were defined on a structural basis, the analysis was necessarily qualitative. He stated that the *incipient* palsa is aggrading and that the *mature palsa* is degrading. Zoltai stated that *young* palsas are usually less than 20 cm in height, whereas *mature* palsas are 100 cm or higher. He also defined palsas by their areal extent, stating that a *young* palsa is generally less than 200 m² while a *mature* palsa may reach several km².

Vegetation is a critical factor in Zoltai's (1972) description of the developmental stages of palsa formation. However, a less qualitative and clearer understanding of the stages of palsa development can be made by using vegetation in conjunction with palsa morphology, surface height above a water level datum, areal extent, active layer thickness and permafrost thickness.



3.4 PALSA DEVELOPMENT: MINEROTROPHIC TOPOGENOUS CARR PEATLAND

The present study develops field criteria to recognize developmental stages of minerotrophic topogenous carr palsas and to differentiate morphologically between these successive stages. The research is based on intensive field work conducted in three separate depressions in the study area. The depressions are occupied by small shoaling lakes surrounded by sedge meadows. During the spring snowmelt period, the meadows accumulate a shallow layer of water. By late summer, soil frost is virtually absent from the meadows except for lenses of permafrost that occur under hummocks. In some instances, contiguous to these meadows uplifted peatland contains a perennially frozen core.

3.4.1 PROCEDURE

The following procedures were used to determine the developmental stages of minerotrophic topogenous carr palsas.

- Twelve transects were surveyed across the depressions of the study area.
- 2. The height of peat hummocks (having a perennially frozen core) above the sedge meadow water level datum (23 July) was determined by a survey rod and tape measure.
- 3. Weekly depth to soil frost measurements were made at nodal points within each plant association using a Hoffer probe.



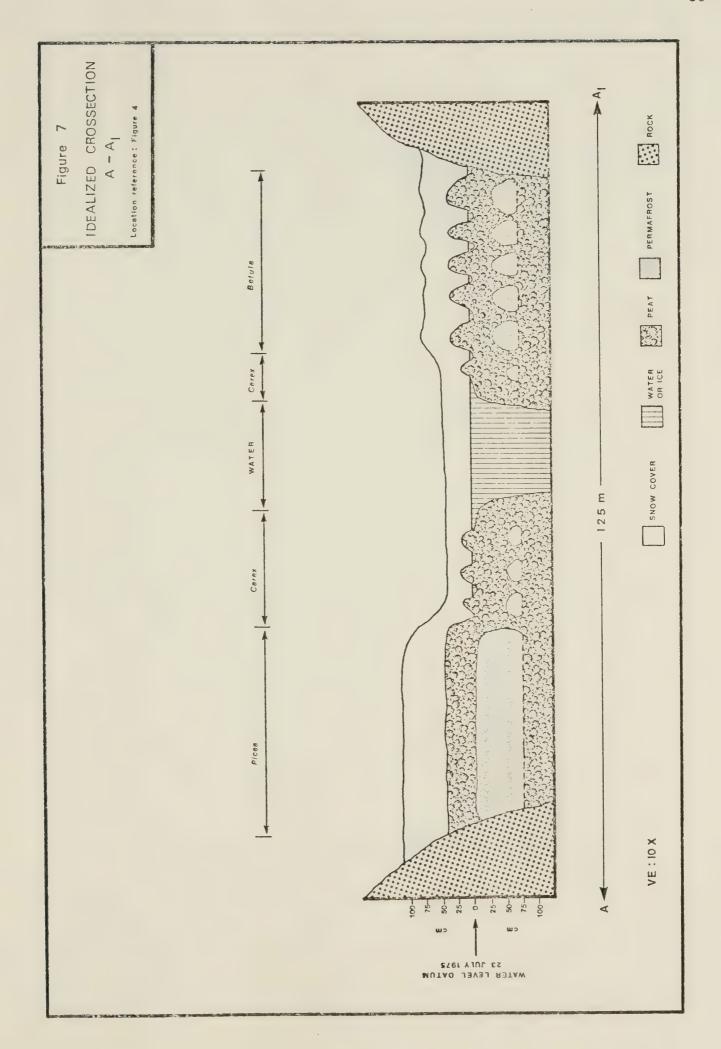
- 4. An Abney level was used to determine the nodal point heights of the *Ledum Cladonia* and *Picea Cladonia* plant associations above the water level datum of the sedge meadow. However, as the height differences were minimal, I found that the instrument did not provide sufficiently accurate measurements.
- 5. Snow depth measurements were made during March, 1975 at nodal points within each plant association using a meter stick.

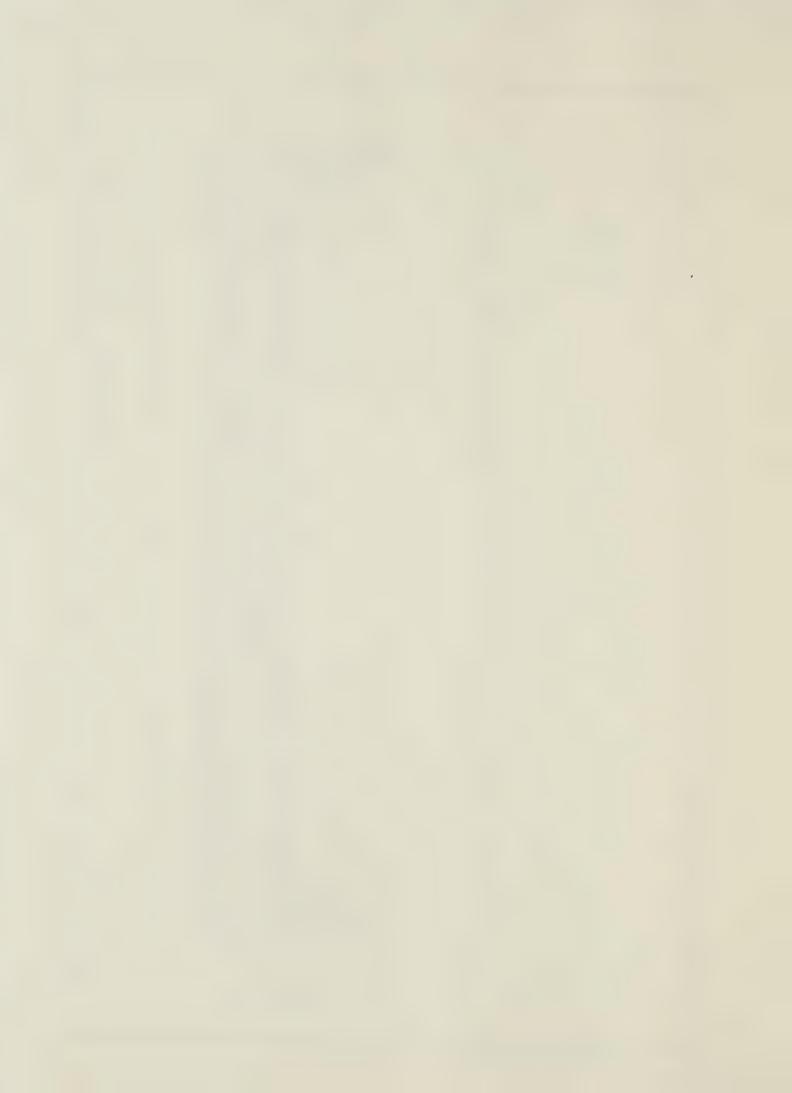
3.4.2 HUMMOCKS

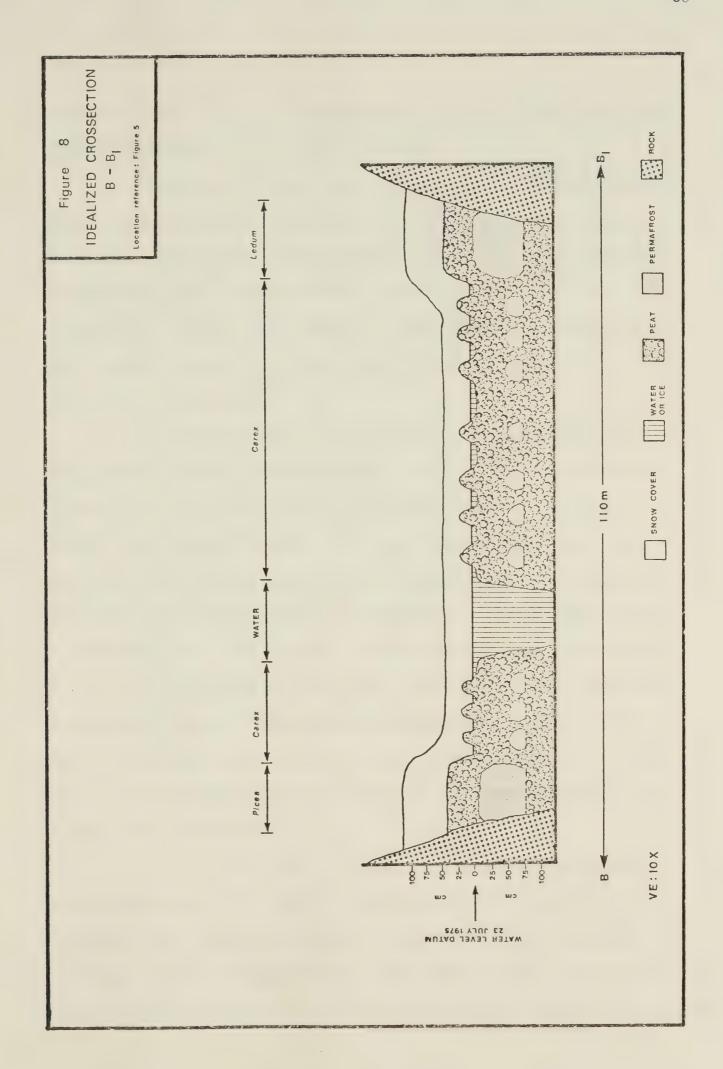
Depressions of the study area are occupied by small shoaling lakes surrounded by sedge meadows of the Carex - Tomenthypnum plant association (Figures 4, 5, 7, 8). During spring and early summer the meadows are flooded to a shallow depth by snowmelt runoff. During particularly wet years, as was the case in 1974, sedge meadows may be flooded all summer. In 1974, seasonal soil frost thawed rapidly because water (acting as a heat source) remained on the meadow, and underlying peat remained saturated allowing rapid conductive heat transfer. Buckman and Brady (1969) estimated that heat passes from water to soil 150 times more rapidly than from air to soil.

With the exception of permafrost lenses under scattered hummocks, soil frost was absent from the *Carex - Tomenthypnum* meadows by 23 July (Figure 6, Appendix 2). These hummocks varied from 14 to 35 cm in height above the 23 July water level datum of the meadow and had an











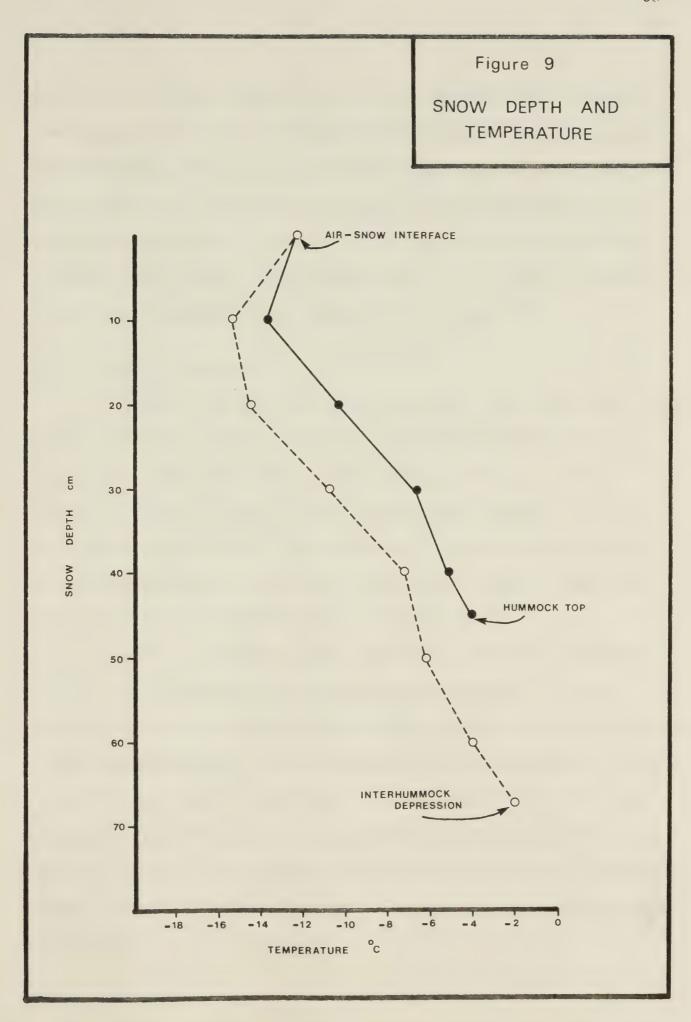
average relief of 24.2 \pm 1.5 cm (Appendix 3). Mean active layer depth within the hummocks was 68 cm on 27 August.

Continguous to the Carex - Tomenthypnum hummocks are other hummock areas dominated by the Betula - Pohlia association (Figures 4, 5, 7). These hummocks varied from 27 to 43 cm in height above the 23 July water level datum and had a mean relief of 34.5 ± 1.0 cm (Appendix 3). On 27 August frost was absent between the hummocks. However, average active layer depth within the hummocks was $64.0 \pm .87$ cm (Figures 6, Appendix 2).

The development of permafrost lenses in certain areas of sedge meadows but not in others appears to be caused by the following conditions. Fall rains, which are followed by low snowfalls in early winter, are important factors. Fall rain causes sedge tussocks and peat hummocks to become wet, which facilitates heat transfer (Buckman and Brady 1969) to the surface by conduction. Freezing temperatures cause further heat loss from the tussocks and hummocks, particularly if there is little snow to retard heat transfer. As snow accumulates during early winter, inter-hummock depressions fill before the hummocks are covered. The thermal resistance created by early and deeper snow between hummocks acts as an effective localized insulating layer against frost penetration.

By March 1975, snow depth was approximately 22 cm deeper between hummocks of the Betula - Pohlia plant association than on the hummock crests (Figure 9, Appendix 4). Concomitantly, the hummock snow-peat interface temperature was approximately 2° C colder than the snow-peat interface temperature of the inter-hummock depression (Figure 9).







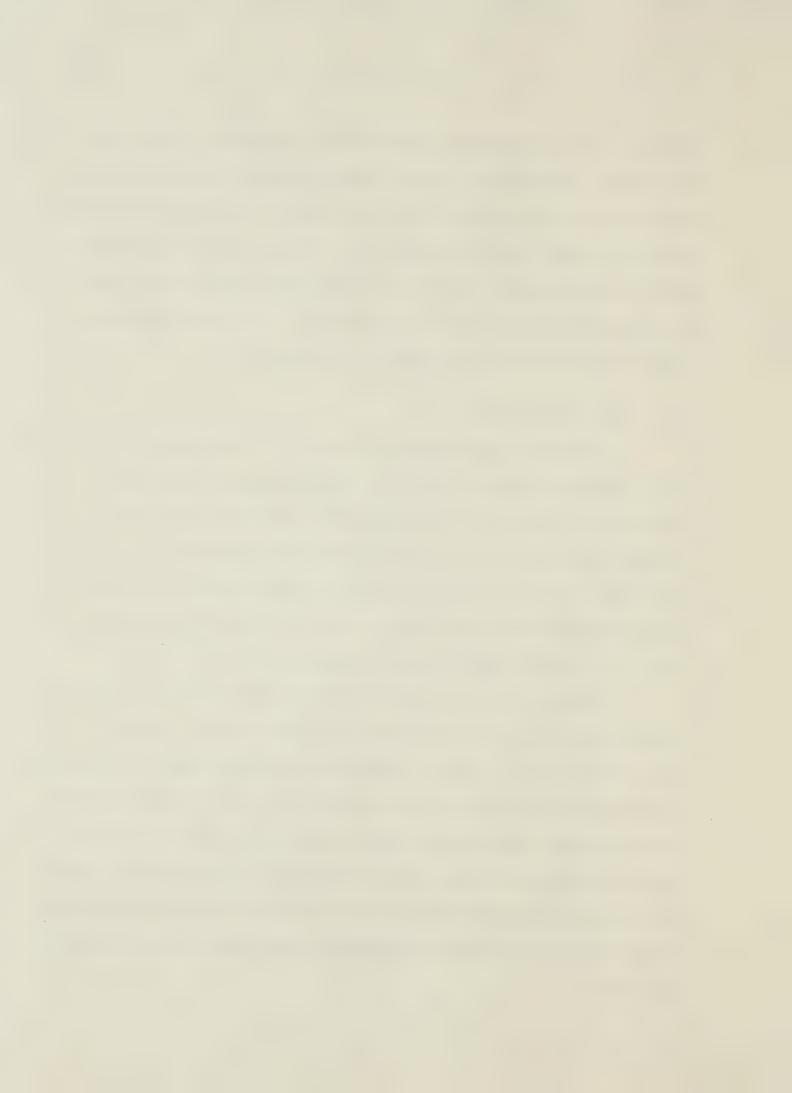
Although similar measurements were not made on hummocks of the <code>Carex</code> - <code>Tomenthypnum</code> association, I assume that the greater snow depth between those hummocks also acted as a localized thermal resistance to heat loss. During the summer, uplifted hummock peat is better drained and becomes drier than meadow peat, therefore its insulating effect is sufficient to reduce heat exchange above the permafrost. As a result, permafrost lenses that form beneath the hummocks are preserved.

3.4.3 UPLIFTED PEATLAND

Proximal to the hummocks of the Carex - Tomenthypnum and Betula - Pohlia plant associations are areas of undulating peatland having a perennially frozen core. These palsas are occupied by the Ledum - Cladonia and Picea - Cladonia plant associations (Figures 4, 5, 7, 8).

On 23 July, the surface of these palsas was approximately 45 cm above the sedge meadow water level datum. Mean active layer thickness was 46.1 + .75 cm on 27 August (Figure 6, Appendix 2).

Palsa development results from slow peat infilling of depressions as well as permafrost aggradation between hummocks. During summer in the <code>Ledum - Cladonia</code> association, the high albedo of <code>Cladonia</code> spp. lichen coupled with the insulating effect of dry uplifted peat is sufficient to reduce heat exchange above the permafrost. Within the <code>Picea - Cladonia</code> association, trees displace the energy exchange surface upward from the air-peat interface and further reduce heat exchange above permafrost. Therefore, an increasing amount of peat remains frozen throughout the summer.



3.4.4 THERMOKARST FEATURES

Shallow, water-filled thermokarst depressions are present on some palsas occupied by the <code>Picea - Cladonia</code> plant association. Active layer thickness beneath one thermokarst pool was 54 cm on 27 August. It appears that these features are initiated when water collects in shallow depressions during periods of high precipitation. The organic material beneath them thus remains saturated, which facilitates heat transfer (Buckman and Brady 1969) by conduction, resulting in thawing of the permafrost.

3.5 SUMMARY

Palsa uplift (development of permafrost) is principally caused by water migration to the freezing plane, volume expansion of water upon freezing (Rapp and Rudberg 1964, Forsgren 1966, Salmi 1968, Lundqvist 1969, Zoltai 1972) and by the buoyancy effect created by frozen organic material overlying saturated unfrozen peat (Salmi 1968, Zoltai 1972). In some peatlands, permafrost and accompanying palsa development is initiated under dense clumps of *Picea mariana* (Zoltai 1972). The present research indicates that in minerotrophic topogenous carr peatland, permafrost genesis and palsa initiation occurs in sedge meadows occupied by the *Carex - Tomenthypnum* plant association. Over time, slow peat infilling of depressions between permafrost hummocks as well as permafrost aggradation creates an uplifted peatland having a perennially frozen core. These palsas are then occupied by the *Ledum - Cladonia* and *Picea - Cladonia* plant associations.



CHAPTER FOUR

SMALL MAMMALS

4.1 INTRODUCTION

Most small mammal research near Great Slave Lake has been centered at the Heart Lake Research Station of The University of Alberta. Similar research in the vicinity of the north arm of Great Slave Lake is limited. As well, no previous work is available that studies the suitability of minerotrophic topogenous carr peatland as small mammal habitat or determines which particular plant associations of this peatland contain the most number of small mammal species.

4.2 PROCEDURE

To understand the interrelations of small mammals and differing palsa habitats the following procedures were used.

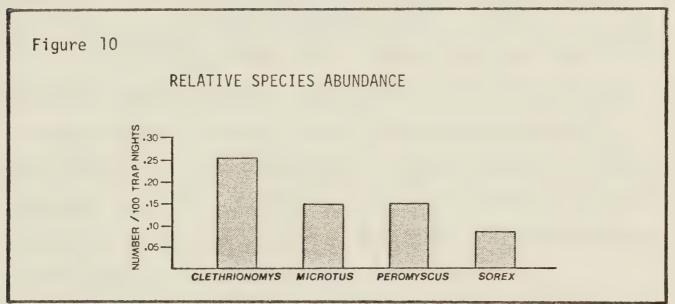
- Eleven snap trap lines were set across the three depressions of the study area. Each line extended
 20 m onto bedrock adjacent to the peatland.
- Trapping stations were spaced at 5 m intervals along the lines. Two traps (baited with peanut butter) were placed at each station.
- 3. Each line was left in position for two 20-day periods (24 June 14 July, 1 August 21 August) and checked daily.



4. Each trapped animal was weighed and length measurements were made of the right hind foot, right ear, tail and body. Skulls of the animals were processed and returned to The University of Alberta where final identification was made by Dr. W. A. Fuller, Department of Zoology.

4.3 SPECIES DISTRIBUTION

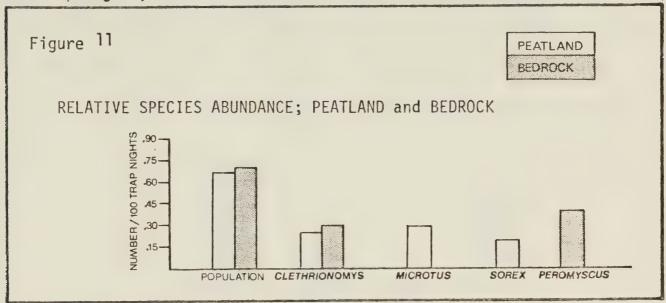
One species of shrew (Sorex cinereus) and three rodents
of the Family Muridae (Clethrionomys rutilus, Peromyscus maniculatus
and Microtus pennsylvanicus) were trapped in the study area. By combining
data from both trapping periods and from all habitats, the relative
abundance of each species was calculated (Figure 10).



Clethrionomys rutilus is clearly the most numerous small
mammal (40 percent of catch) of the study area. Microtus pennsylvanicus
and Peromyscus maniculatus are equally present and comprise approximately
23 percent of the catch. Sorex cinereus represents 14 percent of the catch.



Figure 11 compares the total population and the abundance of each species trapped on the peatland to that of bedrock. Although somewhat higher on bedrock (.68 animals per 100 trap nights), species richness is not significantly greater than peatland (.64 animals per 100 trap nights).



Clethrionomys rutilus has a relatively even distribution in bedrock and peatland habitats. However, the other three species are much more restricted in habitat range. Microtus pennsylvanicus and Sorex cinereus were exclusively trapped in peatland habitats, whereas Peromyscus maniculatus was trapped only in bedrock habitats. As shown by Banfield (1974), Peromyscus maniculatus seldom selects wet habitats, whereas Microtus pennsylvanicus prefers such areas.

Within the peatland, species abundance is greatest in the Betula - Pohlia plant association (Figure 12). Due to a continuous water cover throughout the summer, trapping of small mammals was not conducted in the Carex - Tomenthypnum plant association. However,

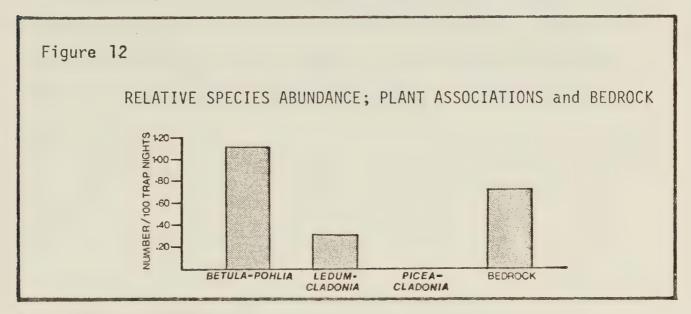


research indicates that at each stage of succession toward the climax Picea - Cladonia plant association, species abundance decreases.

The reasons for this are unknown, but they may be attributed to the increase of Ledum palustre var. decumbers (Fuller, personal communication, 1976). Ledum palustre var. decumbers contains ledol, a

poisonous substance causing cramps and paralysis in humans (Hulten

1968), which may also have an adverse effect on small mammals.



4.4 SPECIES DIVERSITY

Species diversity of the Betula - Pohlia, Ledum - Cladonia and Picea - Cladonia plant associations, as well as the bedrock habitat is calculated using Brillouin's formula (Pielou 1966, Table 11, Appendix 5). Eveness of the plant associations was determined using Pielou's (1966) formula (Table 11).

TABLE 11: SPECIES DIVERSITY AND EVENNESS

	Species Diversity	Evenness
Habitat 1 Betula - Pohlia	.46	.76
Habitat 2 Ledum - Cladonia	.45	.75
Habitat 3 Picea - Cladonia	0	0
Habitat 4 Bedrock	.29	.49



4.5 SUMMARY

Four species of small mammals (Sorex cinereus, Clethrionomys rutilus, Peromyscus maniculatus and Microtus pennsylvanicus) were trapped in the study area. Equally present in the bedrock and peatland habitat, Clethrionomys rutilus represents 40 percent of the catch. However, Microtus pennsylvanicus, Peromyscus maniculatus and Sorex cinereus are much more restricted in habitat range. At each stage of succession toward the climatic climax, species abundance, diversity and evenness decreases.



CHAPTER FIVE

CONCLUSIONS

5.1 INTRODUCTION

This research describes the sequence of environmental conditions that contribute to the development of minerotrophic topogenous carr peatland ecosystems near Yellowknife, Northwest Territories. The study indicates that permafrost aggradation in organic terrain contributes to the development of distinct physical environments within which characteristic plant associations evolve. Concurrently, differing plant associations provide habitats of varying suitability for small mammals. Although the natural environment does not function in compartments (either in time or space), sectioning is valuable in some instances in reconstructing environmental time sequences (Gill 1972). Therefore, I hypothesize that three stages represent the interaction between minerotrophic topogenous carr palsa development and the formation of small mammal habitat.

5.2 STAGE ONE

Minerotrophic topogenous carr palsa formation initially takes place in Carex - Tomenthypnum meadows where individual peat hummocks develop a perennially frozen core. Snowmelt runoff floods the meadows to a shallow depth during spring and early summer. During particularly wet summers Carex - Tomenthypnum meadows may be flooded continuously.



With the exception of permafrost lenses under hummocks, seasonal soil frost is usually absent from the sedge meadows by August.

Although small mammal trapping was not conducted in the Carex - Tomenthypnum meadows of the present research area, I assume that this plant association has low suitability as small mammal habitat because of its wetness. In particular, Peromyscus maniculatus seldom occurs in wet habitats, Clethrionomys rutilus prefers areas covered by shrubby growths of willow and birch, and Sorex cinereus inhabits moist shrubby areas with a duff ground cover (Banfield 1974).

5.3 STAGE TWO

Over the years, organic accretion and permafrost aggradation cause the individual peat hummocks of the Carex - Tomenthypnum meadows to become larger and more numerous. The surface of these larger hummocks is drier than that of the Carex - Tomenthypnum hummocks, enabling the Betula - Pohlia plant association to become dominant.

Small mammal trapping indicates that species abundance is greatest in the Betula - Pohlia plant association. Therefore, I assume that areas dominated by this association provide the most suitable small mammal habitat of the peatland.

5.4 STAGE THREE

Slow peat infilling of depressions between hummocks as well as permafrost aggradation results in an uplifted palsa having a perennially frozen core. The uplifted palsa surface is drier than



the peat surface of the <code>Betula - Pohlia</code> hummocks which results in the dominance of the <code>Ledum - Cladonia</code> plant association. Over the years, sere development continues and the <code>Ledum - Cladonia</code> plant association is replaced by the <code>Picea - Cladonia</code> association.

Small mammal trapping indicates that species abundance in the <code>Ledum - Cladonia</code> association is less than in the <code>Betula - Pohlia</code> association but higher than in the <code>Picea -Cladonia</code> association. Consequently, I assume that the habitat suitability of minerotrophic topogenous carr peatlands is greatest in the <code>Betula - Pohlia</code> plan association and decreases toward the climatic climax <code>Picea - Cladonia</code> association.

5.5 SUMMARY

Peatland of the present research area is occupied by four separate plant associations. Each plant association possesses a discrete set of environmental conditions which are reflected by three stages of succession to a climatic climax. The differing plant associations provide habitats of varying suitability for small mammals.



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APPENDIX 1

Relevent Terms



For purposes of the present research, the following definitions are adhered to (Mörnsjö 1971; Whittaker 1975):

OMBROTROPHIC Communities that receive nutrients

from the atmosphere.

MINEROTROPHIC Communities that receive nutrients

from both the atmosphere and mineral soil.

TOPOGENOUS Minerotrophic peatlands that develop in

basins containing stagnant ground water.

CARR Topogenous peatlands that are influenced by

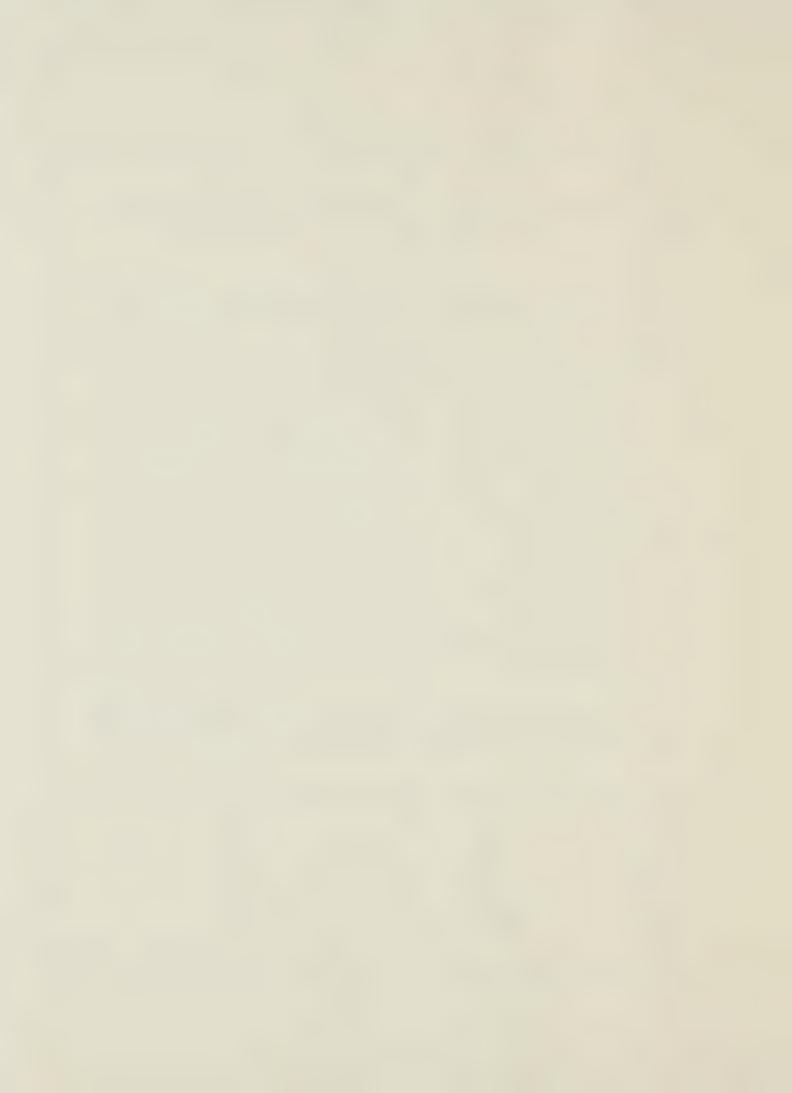
oscillating ground water levels.



APPENDIX 2 Depth To Frost Measurements



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ı×	40.3	65.8	6.89	70.0	72.5	79.7	82.3		20.6	23.1	26.0	29.0	32.7	37.3	41.3	44.5	48.0	52.6	58.2	6.19	64.0	
	40	67	89	7.0	73	81	ŧ		23	27	29	31	35	41	45	48	20	55	62	64	99	
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DATE	8 June	18 June	23 June	2 July	9 July	16 July	23 July	30 July	8 June	13 June	18 June	23 June	2 July	9 July	16 July	23 July	30 July	6 August	13 August	20 August	27 August	
PLANT ASSOCIATION				Carex -	Tomenthypnum										Betula - Pòhlia							



X	15.5 .40	e	. 7.	25.7 .82	30.1 .81	31.5 .72	33.5 .87	37.1 1.01	38.9 .72	40.7	43.4 .93	44.7 .68	46.1 .75	17.3 .51	20.0 .66	22.9 1.19	24.5 1.31	28.5 .98	31.2 1.17	32.8 1.38	35.6 1.41	39.0 1.97	40.4 1.73	43.0 1.96	44.1 1.82	45.8 1.85
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DATE	8 June	13 June	18 June	25 June	2 July	9 July	16 July	23 July	30 July	6 August	13 August	20 August	27 August	8 June	13 June	18 June	23 June	2 July	9 July	16 July	23 July	30 July	6 August	13 August	20 August	27 August
PLANT ASSOCIATION						•	Ledum - Cladonia												Picea - Cladonia							



APPENDIX 3

Hummock Heights Above Water Level Datum 23 July 1975



PLANT ASSOCIATION			工	UMMO	CK H	HUMMOCK HEIGHTS (cm)	TS (cm)			ı×	se.
Carex - Tomenthypnum	20	16	16 18 19 27	19	27	29	25	32	30	26		
	21	17	17 19	29	33	26	16	31	35	14	24.15 1.46	1.46
Betula - Pohlia	28	31	29	27	32	2 40 4	43	36	37	41	1	4
	29	32	34	35	40	39	33	36	37	33	34.50 1.04	1.04



APPENDIX 4

Snow Depth Measurements 25 March 1975



NO THIS CO.	٠		,		84		,			9	:	:	:	:	:	:	:			ŀ	١,	
NOUNE POSTITON	C 2 1	2	2	+	n	0	0		ת	10 11 12 13 14 15 16 17 18 19	=	71	2	4	2	9	_	20		20	×	se;
ASSOCIATION							SN	OM D	SNGW DEPTH (cm	(cm												
Picea - Cladonia (apt)	77	77 77 75	75	68	11	78	64	12	99	2	88	63	63	75	20	65	81	20	99	68	70.3	.1.66
Picea - Cladonla (qamaniq)	62	62 64 63	63	52	52	19	56	57	67	51	22	. 09	63	53	09	65	57	63	65	09	59.7	1.02
Ledum - Cladonia	70	28	99	26	64	59	64	58	52	54	63	20	20	7	63	69	.72	70	64	69	63.5	1.47
Betula - Pohlia (hummocks)	46	46 47 48	48	48	48	49	46	48	42	40	4	48	46	48	46	46	48	40	45	41	45.6	68
Betula - Pohlia (Interhummock depressions)		66 63	99	29	68	70	29	7	70	72	63	64	69	63	17	99	68	29	70	20	67.8	.59
Carex - Tomenthyprum	20	20	50 49	49	49	20	48	48	48	49	45	45	46	45	43	46	46	45	49	20	47.6	.47



APPENDIX 5
Species Diversity And Evenness



Information content is often used as a measure of the diversity of many-species biological collections (Pielou 1966).

Diversity is defined as the degree of uncertainty attached to the specific identity of any randomly selected individual (Pielou 1966). The greater the number of species and the more nearly equal their proportions, the greater the diversity (Pielou 1966).

For the present study, species diversity was measured by Brillouin's formula:

S.D. =
$$\frac{1}{N} \log \frac{N!}{N_1! N_2! \dots N_s!}$$

where:

N is the total number of individuals

s is the total number of species

 N_1 is the total number of individuals in the 1th species

The diversity of a collection of species is maximum when the individuals are distributed as evenly as possible among the species.

The ratio of the observed diversity to the maximum possible diversity is a measure of the evenness with which the individuals are divided among the species.

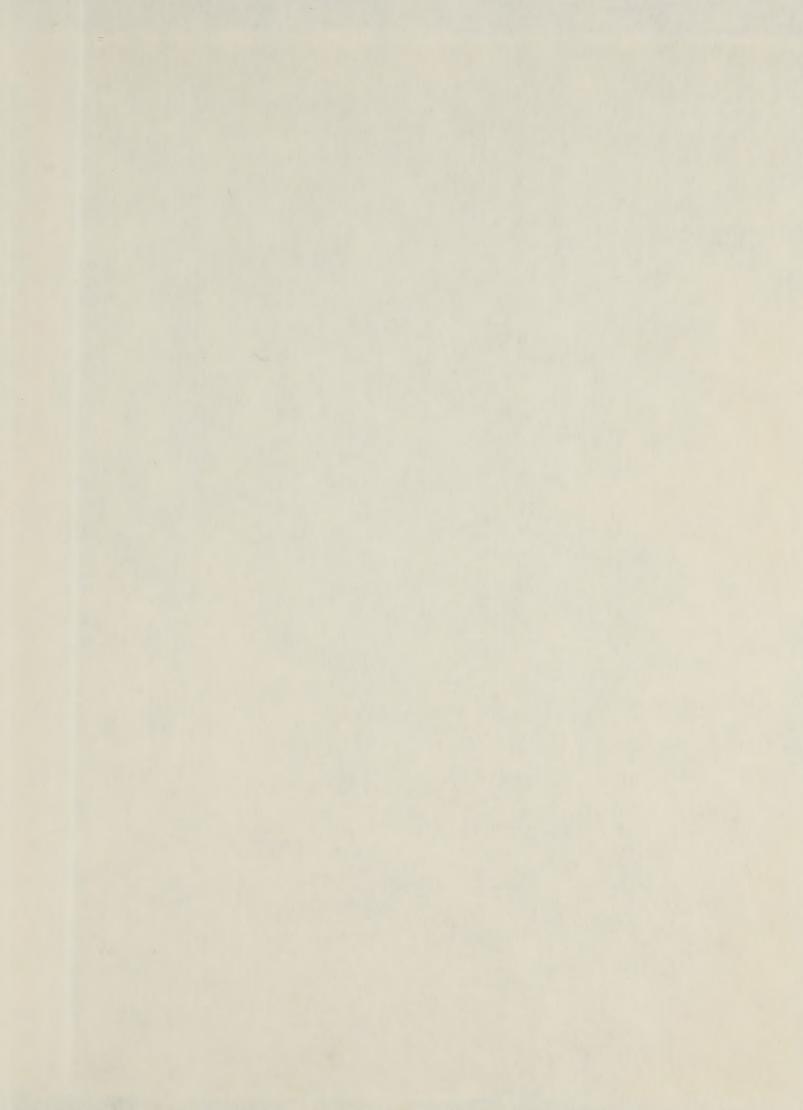












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